



12th DECIGO Workshop
東京大学理学部4号館1220室
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今後の衛星重力ミッションの動向

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話 題

- ・ 衛星重力ミッションの概要
- ・ これまでの成果
 - 時間変動地球重力場(GRACE)
 - 静的(平均)地球重力場(GOCE)
- ・ 将来ミッション

衛星重力ミッション

CHAMP (2000.07.15)- (2010.09.19)

- CHALLENGING Mini-satellite Payload for geophysical research applications
- H-L Satellite-to-Satellite Tracking

GRACE (2002.03.17)- (2017 ??)

- Gravity Recovery And Climate Experiment
- L-L Satellite-to-Satellite Tracking (Microwave Link)

GOCE (2009.03.17) - (2013.10.21)

- Gravity field and steady-state Ocean Circulation Explorer
- Gravity Gradiometer

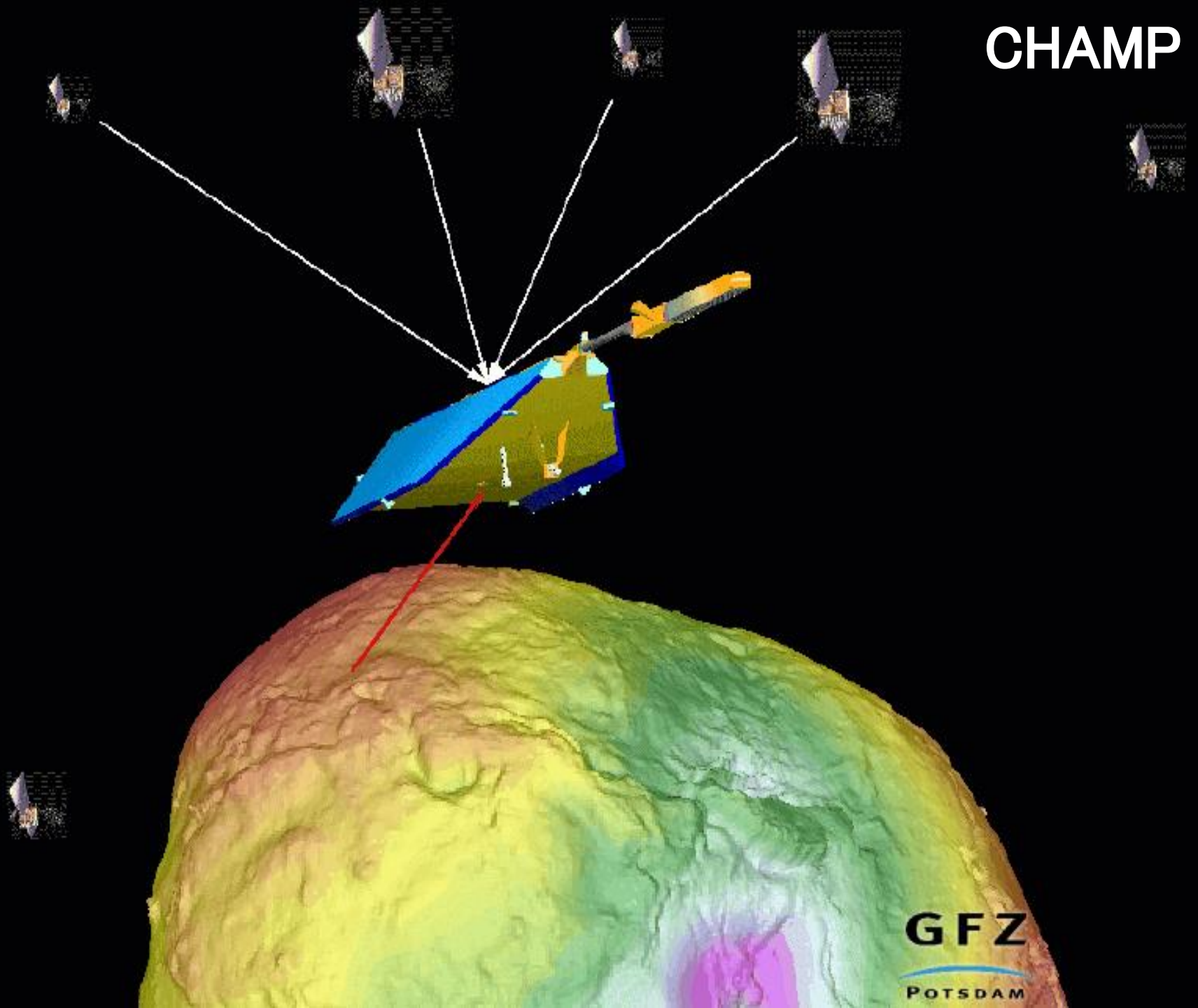
Swarm(2013.11.14 scheduled)

- ESA's magnetic field mission
- H-L Satellite-to-Satellite Tracking (CHAMP like 3 satellites)

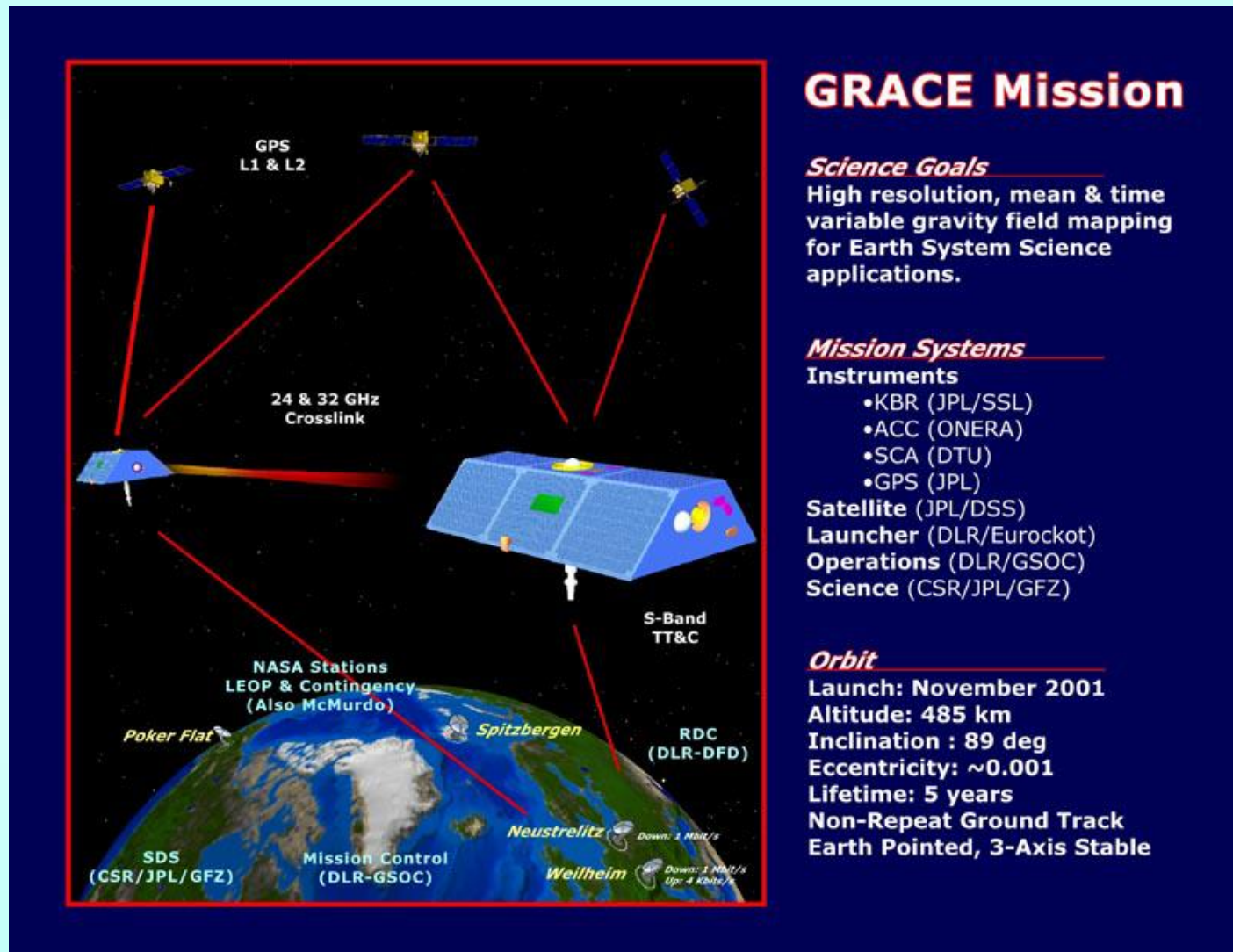
GRACE-FO (2017), GRACE-II (202?)

- GRACE GAP Filler (Microwave Radar)
- L-L Satellite-to-Satellite Tracking (Laser Link)

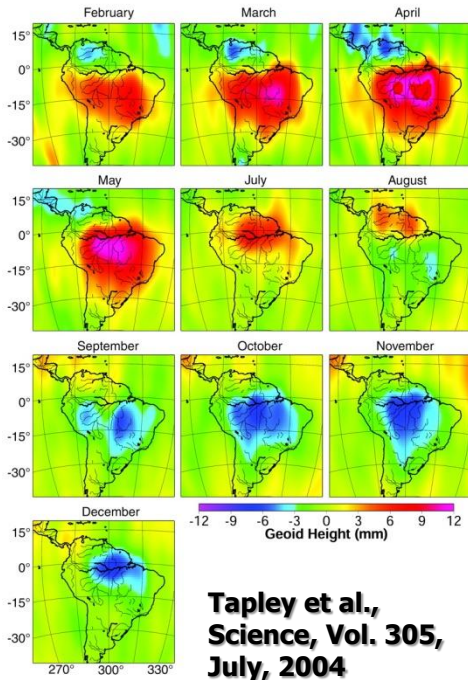
CHAMP



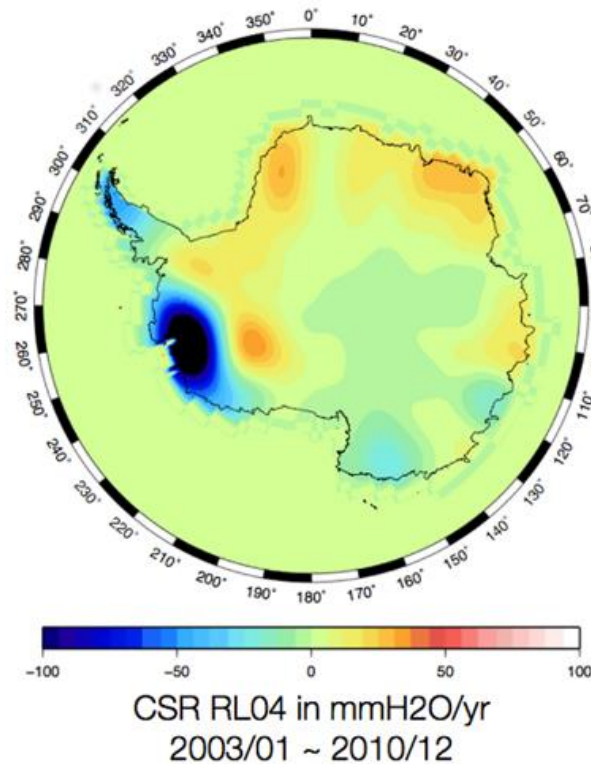
GRACE Mission



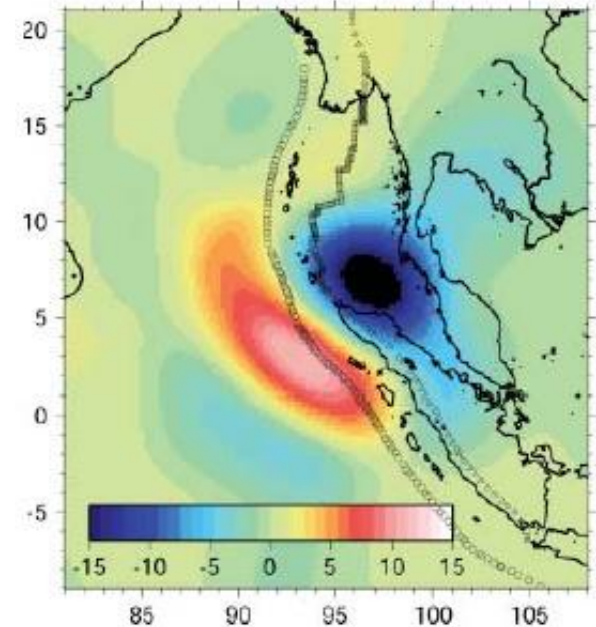
GRACE Applications



**Landwater
Changes**



**Ice Sheet Mass
Changes**



Earthquakes

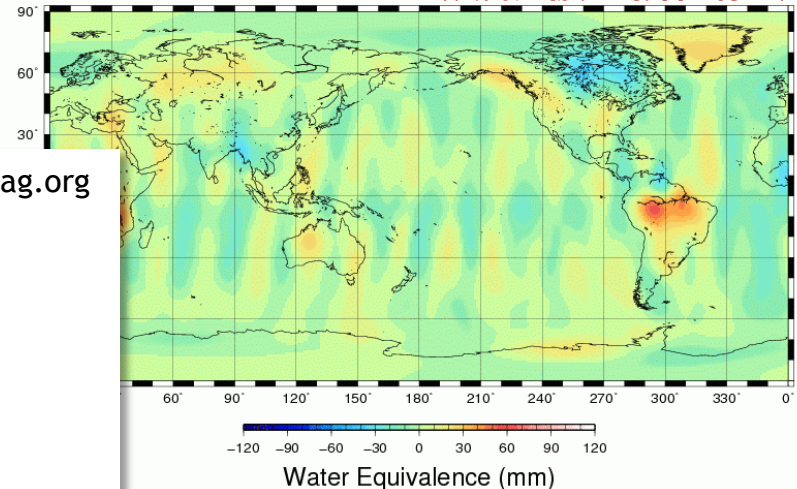
現行の衛星重力ミッション～ GRACE (2002～)



GRACEデータから求めた地表面密度の時間変化

(大気、潮汐の影響は除く)

d001-2002-102



REPORTS

24 MARCH 2006 VOL 311 SCIENCE www.sciencemag.org

Measurements of Time-Variable Gravity Show Mass Loss in Antarctica

Isabella Velicogna^{1,2*} and John Wahr^{1*}

Using measurements of time-variable gravity from the Gravity Recovery and Climate Experiment satellites, we determined mass variations of the Antarctic ice sheet during 2002–2005. We found that the mass of the ice sheet decreased significantly, at a rate of 152 ± 80 cubic kilometers of ice per year, which is equivalent to 0.4 ± 0.2 millimeters of global sea-level rise per year. Most of this mass loss came from the West Antarctic Ice Sheet.

南極氷床の減少を初めて明らかに

2002～2005年のGRACEデータから、南極全体で $152 \pm 80 \text{ km}^3/\text{yr}$ で質量が減少していることを報告

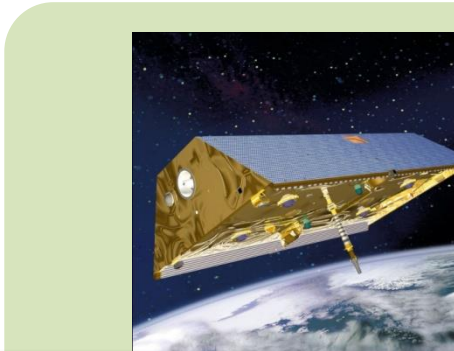
- ✓ 地上観測が困難な環境での大陸全体での氷床質量の減少量を
→ 衛星観測の有効性

精度： 10^{-8} m/s^2 の精度の全球解を、
1カ月間隔

→ 地表での水の変化に換算して数cm



現行の衛星重力



24 MARCH 2006 VOL 311 SCII

REPORTS

Measurements of Time-Varying Gravity Show Mass Loss in the Antarctic Ice Sheet

Isabella Velicogna^{1,2*} and John Wahr^{1*}

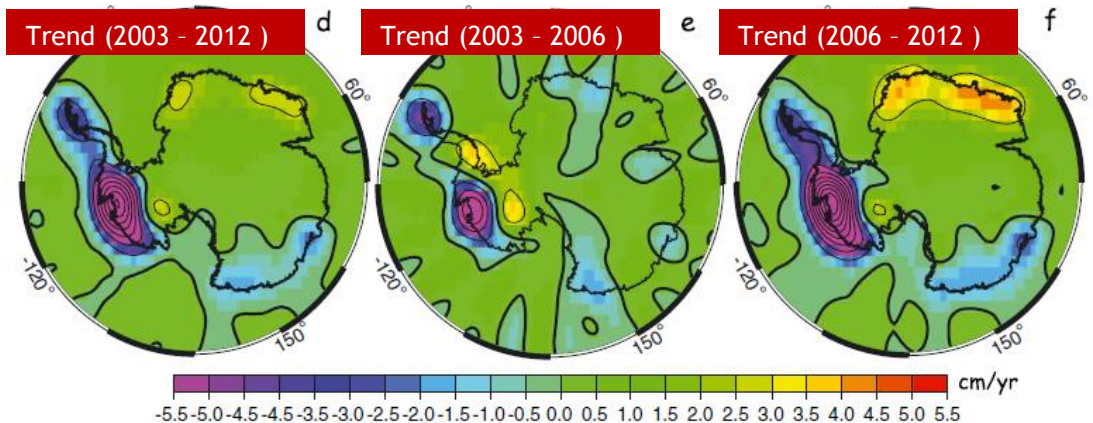
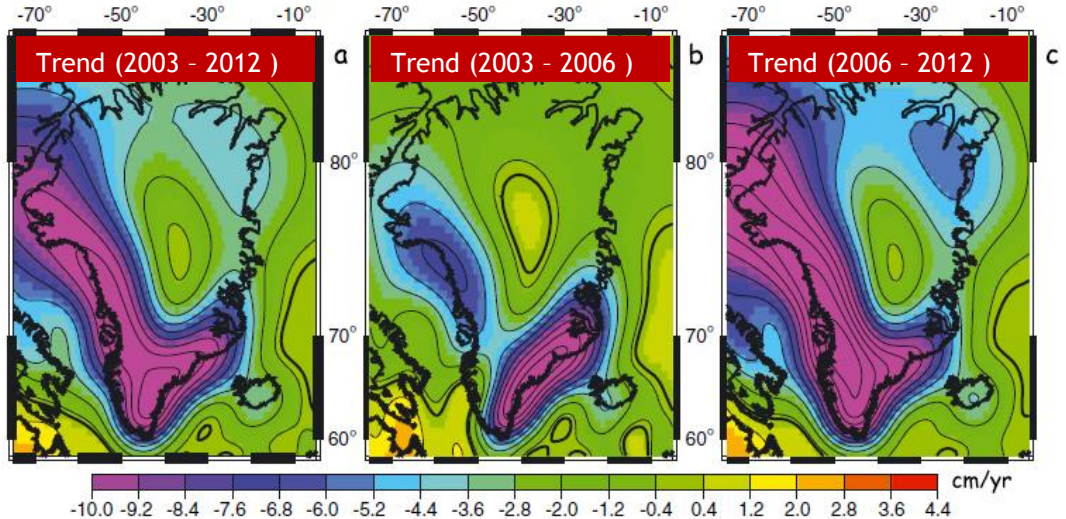
Using measurements of time-variable gravity from the Gravity Recovery and Climate Experiment (GRACE) satellites, we determined mass variations of the Antarctic ice sheet and found that the mass of the ice sheet decreased significantly, at a rate of 150 Gt per year, which is equivalent to 0.4 ± 0.2 millimeters of global sea-level rise per year from the West Antarctic Ice Sheet.

南極氷床の減少を初めて

2002～2005年のGRACEデータから、 $2 \pm 80 \text{ km}^3/\text{yr}$ で質量が減少していることが明らかになった。

- ✓ 地上観測が困難な環境での大陸全体での観測が可能
- 衛星観測の有効性

GEOPHYSICAL RESEARCH LETTERS, VOL. 40, 3055-3063, doi:10.1002/grl.50527, 2013



近年の氷床融解の加速

大陸で平均すると、質量減少が加速

地域別には、氷床融解が加速する地域と

積雪により質量が増加した地域の両方がある

現行の衛星重力



24 MARCH 2006 VOL 311 S

REPORTS

Measurements of Time-Varying Gravity Show Mass Loss

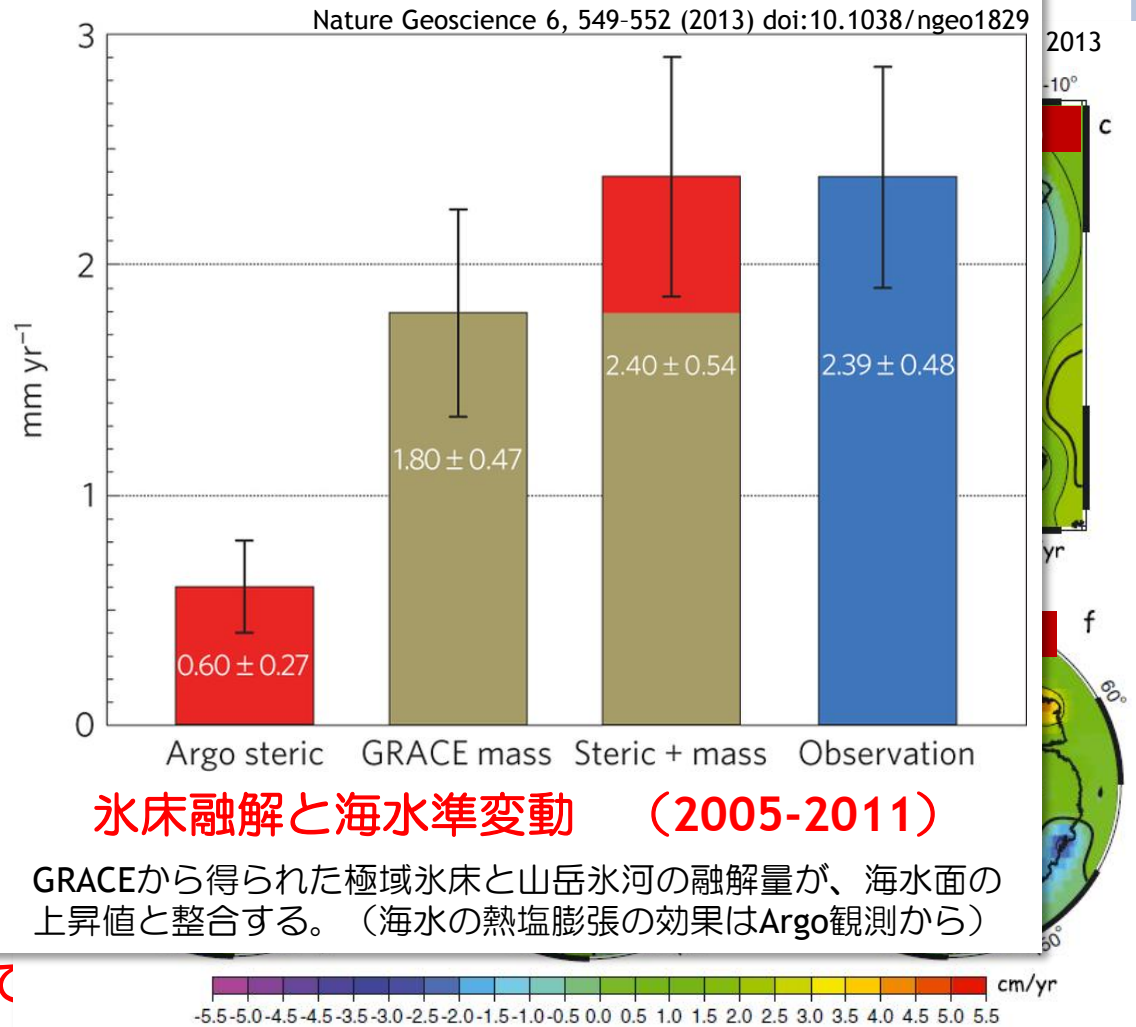
Isabella Velicogna^{1,2*} and John Wahr^{1*}

Using measurements of time-variable gravity from the Gravity Recovery and Climate Experiment (GRACE) satellites, we determined mass variations of the Antarctic ice sheet that the mass of the ice sheet decreased significantly, at a rate of 1.80 ± 0.47 mm yr⁻¹ per year, which is equivalent to 0.4 ± 0.2 millimeters of global sea mass loss came from the West Antarctic Ice Sheet.

南極氷床の減少を初めて

2002~2005年のGRACEデータ:
2 ± 80 km³/yrで質量が減少していること

- ✓ 地上観測が困難な環境での大陸全体での
→ 衛星観測の有効性



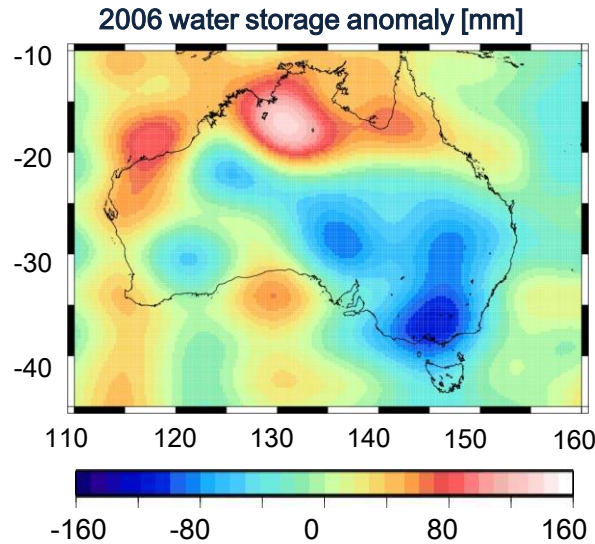
近年の氷床融解の加速

- 大陸で平均すると、質量減少が加速
- 地域別には、氷床融解が加速する地域と
積雪により質量が増加した地域の両方がある

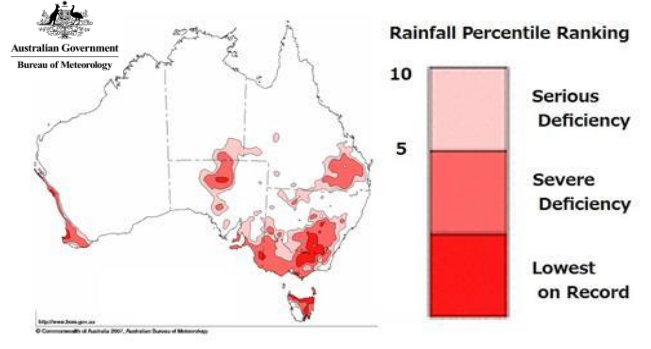
現行の衛星リモートセンシング (GRACE) による

2006年のオーストラリア大旱魃

Hasegawa *et al.*, *EGU General Assembly 2010*



Historic drought struck in 2006 (2006 Drought statement)



Vol 460/20 August 2009 | doi:10.1038/nature08238

Satellite-based estimates of groundwater depletion in India

Matthew Rodell¹, Isabella Velicogna^{2,3,4} & James S. Famiglietti²

インド北部での灌漑による地下水減少

インド北部の地下水が、灌漑により、2002-2008間で $17.7 \pm 4.5 \text{ km}^3/\text{yr}$ 減少

→ 地上モニタリングが困難な地下水変化の衛星観測に成功

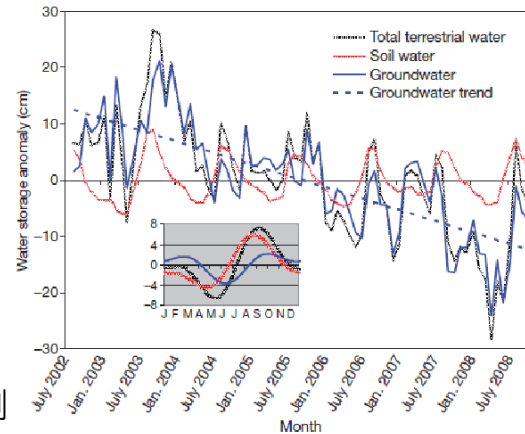


Figure 3 | Monthly time series of water storage anomalies in northwestern India. Monthly time series of anomalies of GRACE-derived total TWS, modelled soil-water storage and estimated groundwater storage, averaged over Rajasthan, Punjab and Haryana, plotted as equivalent heights of water in centimetres. Also shown is the best-fit linear groundwater trend. Inset, mean seasonal cycle of each variable.

全球解を、
に換算して数cm

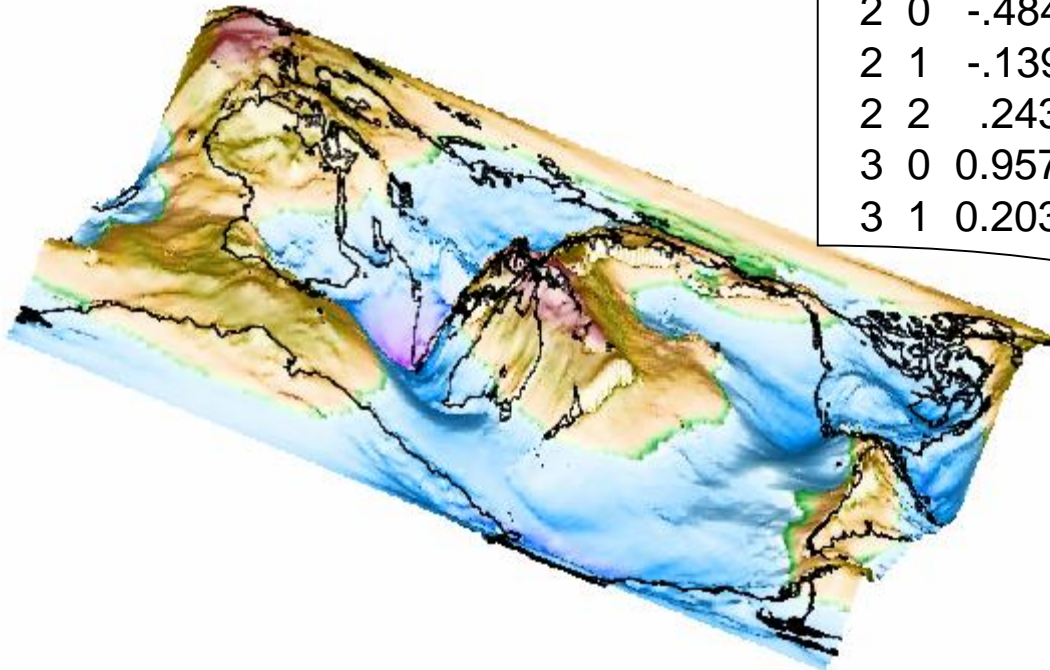


文
言
三
基
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応

重力場データ

$$V = \frac{GM}{R} \sum_{l=0}^L \sum_{m=0}^l \left(\frac{R}{r}\right)^{l+1} \bar{P}_{lm}(\sin \phi) (\bar{C}_{lm} \cos(m\lambda) + \bar{S}_{lm} \sin(m\lambda))$$

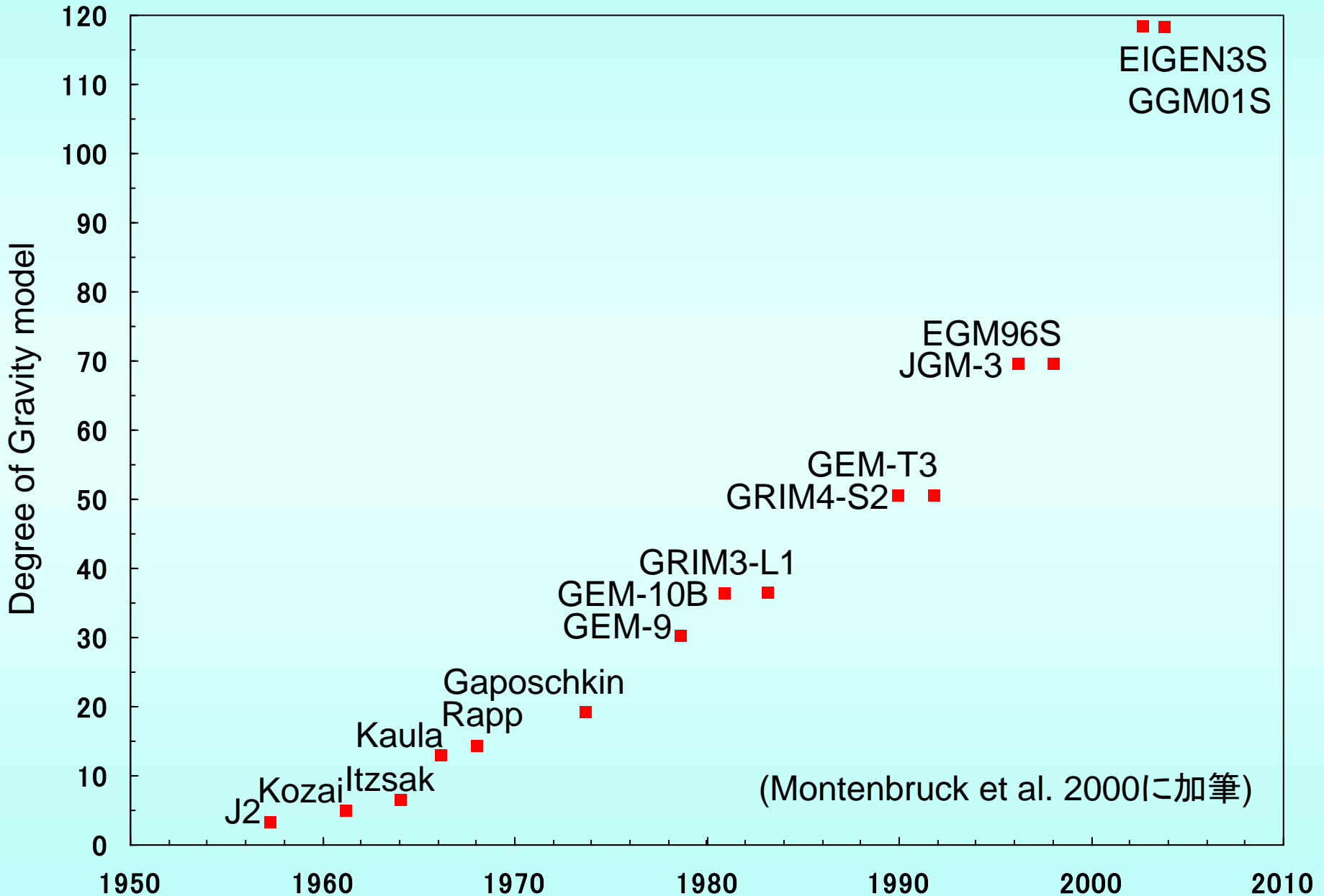
CHAMP
EIGEN-2 model



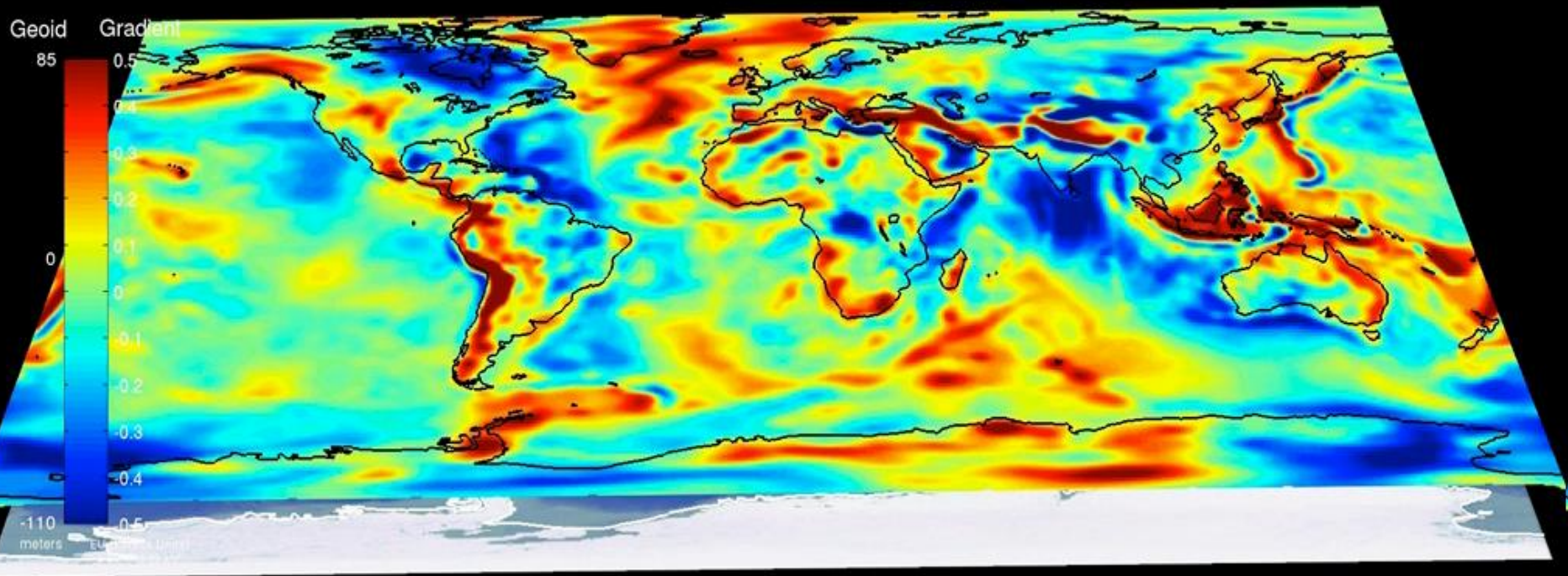
l	m	C_{lm}	S_{lm}
0	0	0.1000000000000D+01	0.0000000000000D+00
1	0	0.0000000000000D+00	0.0000000000000D+00
1	1	0.0000000000000D+00	0.0000000000000D+00
2	0	-.484165815935D-03	0.0000000000000D+00
2	1	-.139274044771D-09	0.141275625216D-08
2	2	.243930491835D-05	-.140027520233D-05
3	0	0.957730216460D-06	0.0000000000000D+00
3	1	0.203054457836D-05	0.248146553306D-06
			⋮

Up to Degree 120 / Order 120

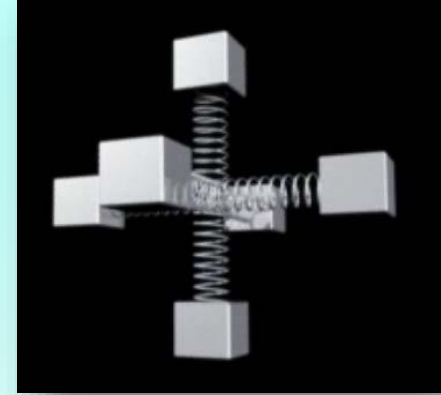
衛星による地球重力場モデルの発展



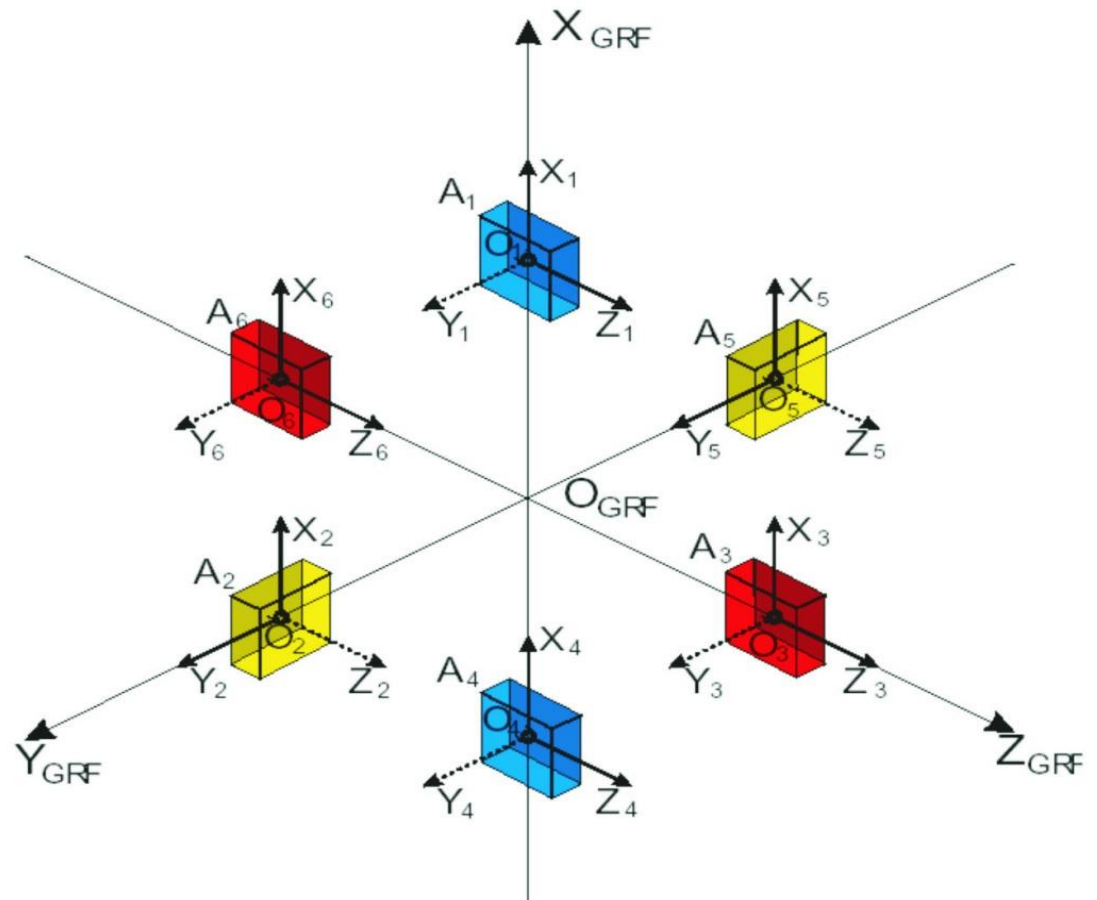
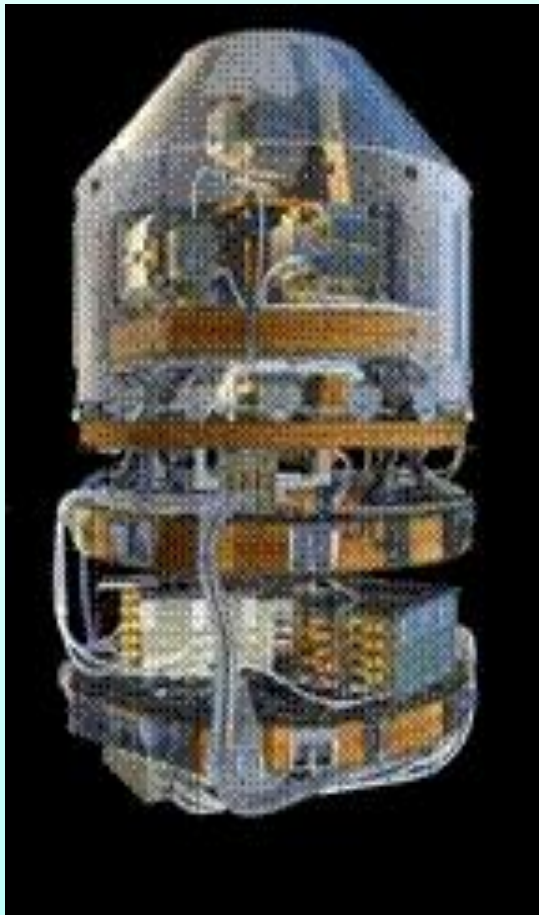
GOCE



EGG

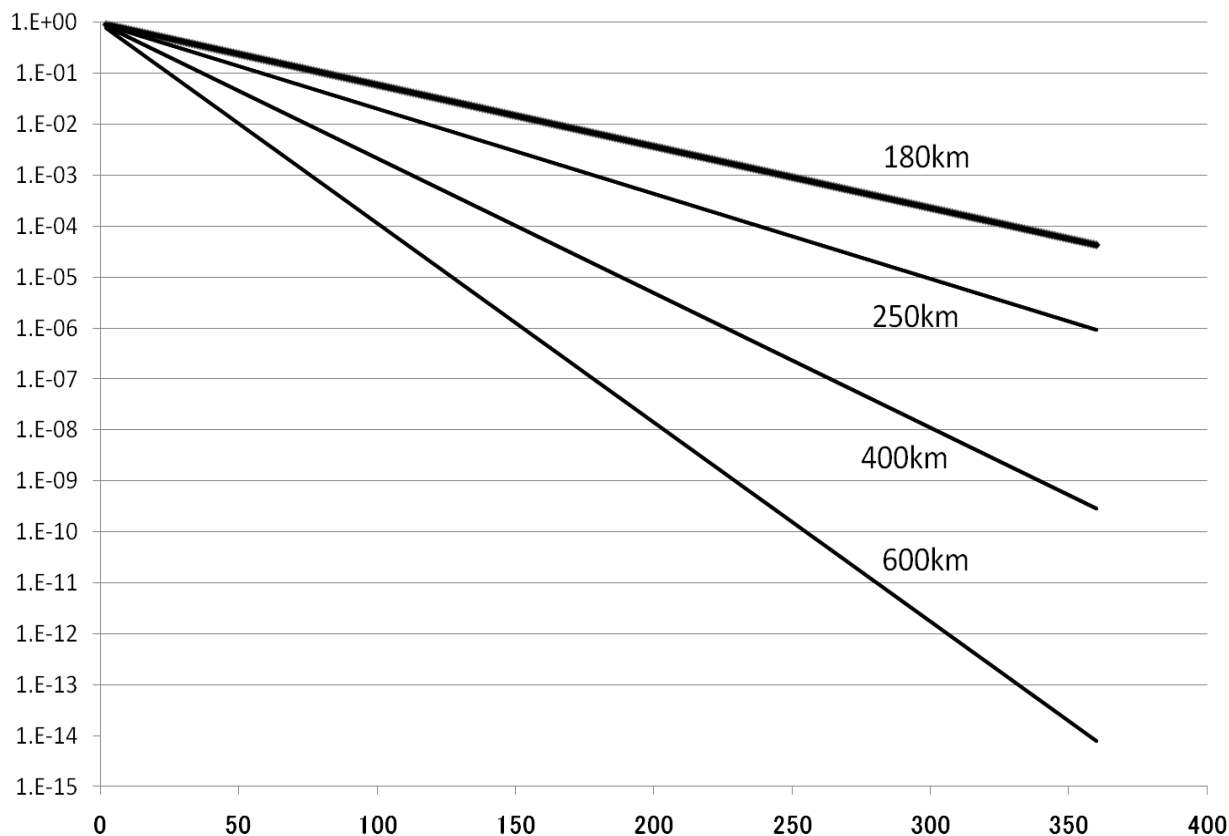


$$[W_{ij}] = \text{grad} \mathbf{g} = \text{grad}(\text{grad}W) = \begin{pmatrix} W_{xx} & W_{xy} & W_{xz} \\ W_{yx} & W_{yy} & W_{yz} \\ W_{zx} & W_{zy} & W_{zz} \end{pmatrix}$$



軌道高度(h)による重力場(V)の減衰

$$V = \frac{GM}{a+h} \sum_{l=0}^{\infty} \sum_{m=0}^l \left(\frac{a}{a+h} \right)^l P_{lm}(\sin\phi) (C_{lm} \cos(m\lambda) + S_{lm} \sin(m\lambda))$$



軌道高度による重力場決定精度

Figure 2.3: 200 km Altitude

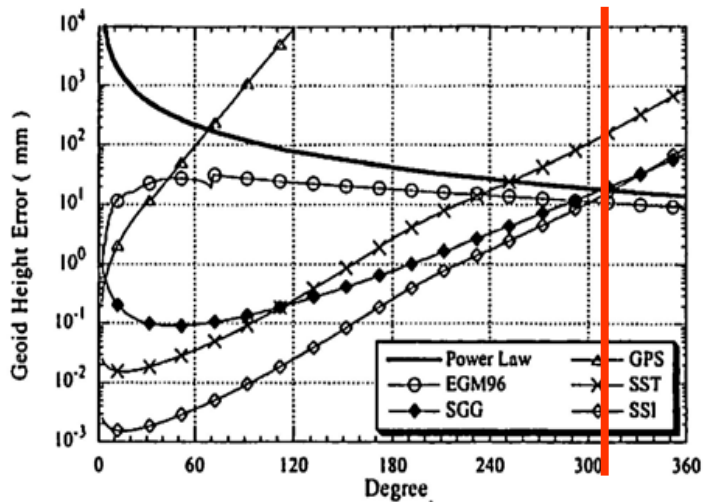


Figure 2.4: 300 km Altitude

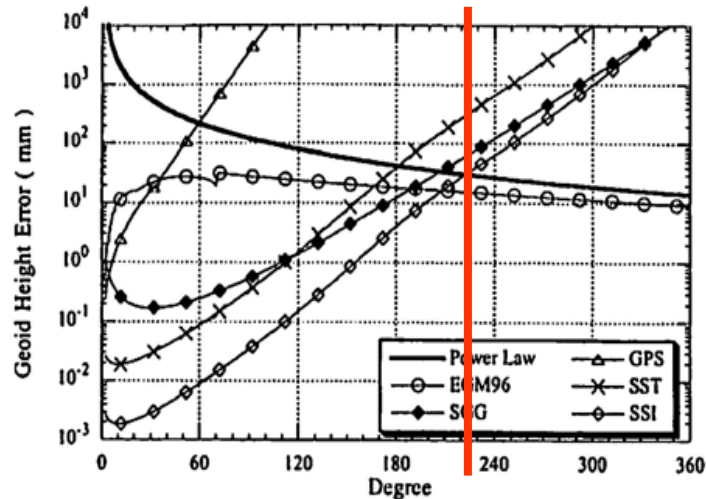


Figure 2.5: 400 km Altitude

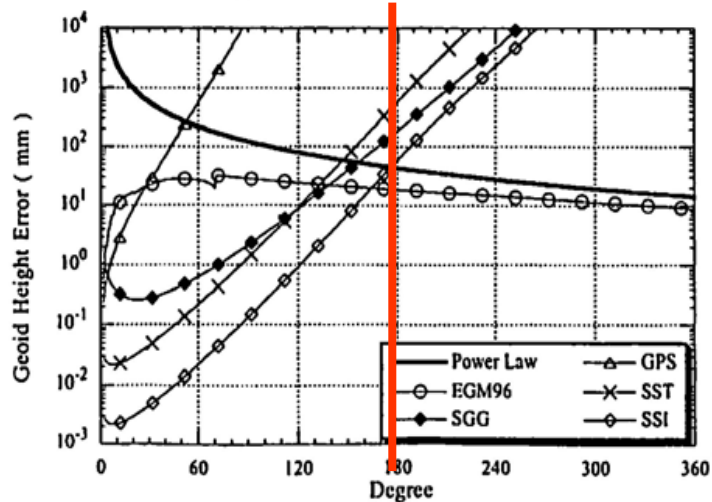
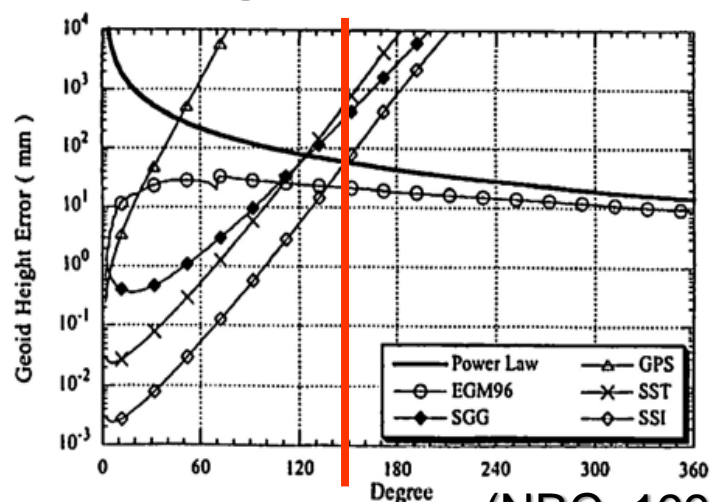


Figure 2.6: 500 km Altitude



地球物理学的应用

Table 3-1: Static gravity field, scientific requirements in preparation for GOCE, from: Rummel (2005).

Application		Accuracy		Spatial resolution
		Geoid [cm]	Gravity [mGal]	Half wavelength: D [km]
Solid Earth	Lithosphere/upper mantle density		1-2	100
	Continental lithosphere	Sedimentary	1-2	50-100
		Basins rifts	1-2	20-100
		Tectonic motions	1-2	100-500
	Seismic hazards		1	100
Ocean lithosphere/asthenosphere		0.5	100-200	
Oceanography	Short scale	1-2		100
		0.2		200
	Basin scale	-0.1		1000
Ice sheets	Rock basement		1-5	50-100
	Ice vertical movements	2		100-1000
Geodesy	Levelling by GPS	1		100-1000
	Unified height system	1		100-20000
	INS		~1-5	100-1000
	Orbits		~1-3	100-1000
Sea level change		Many of the above applications, with their specific requirements, are relevant to studies of sea level change.		

海面形状と海流

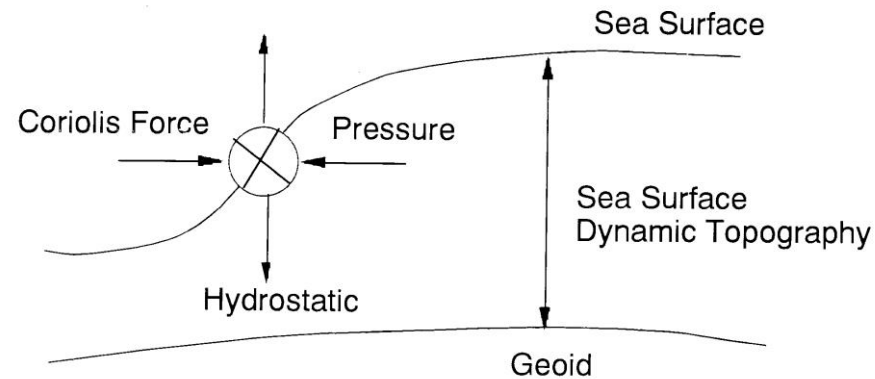
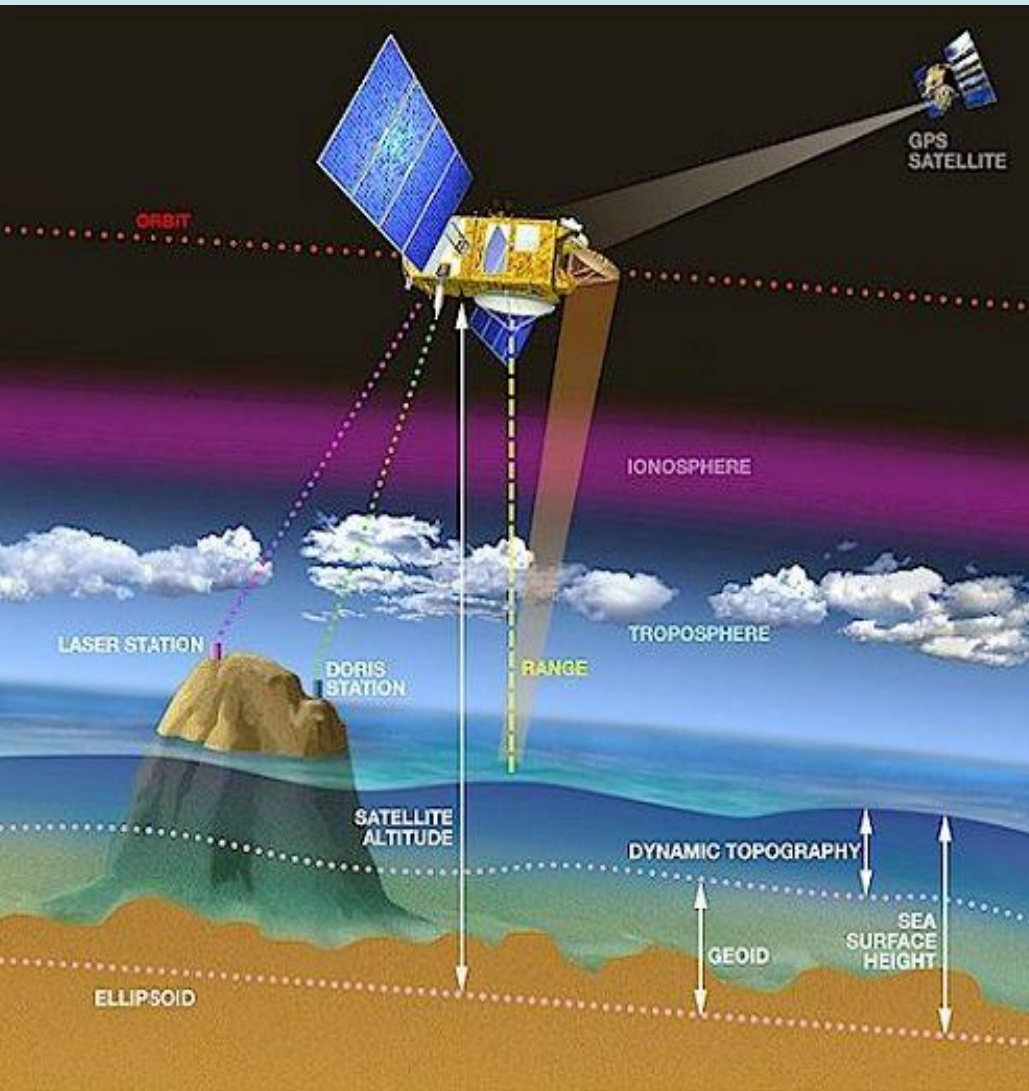
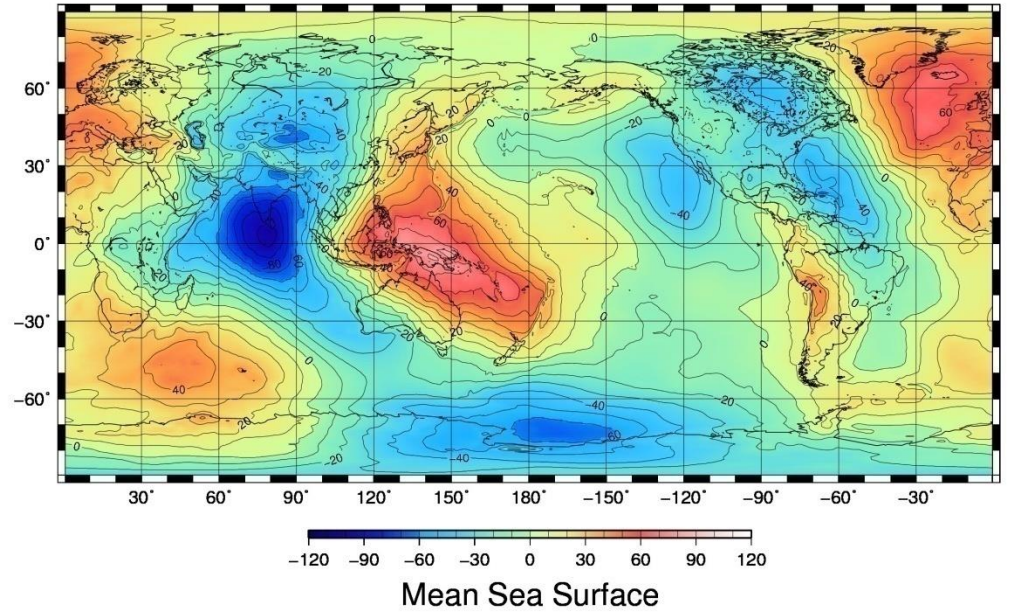


Fig. 2. A schematic view of the sea surface dynamic topography which is caused by geostrophic balance between ocean current and the slope of the sea surface.

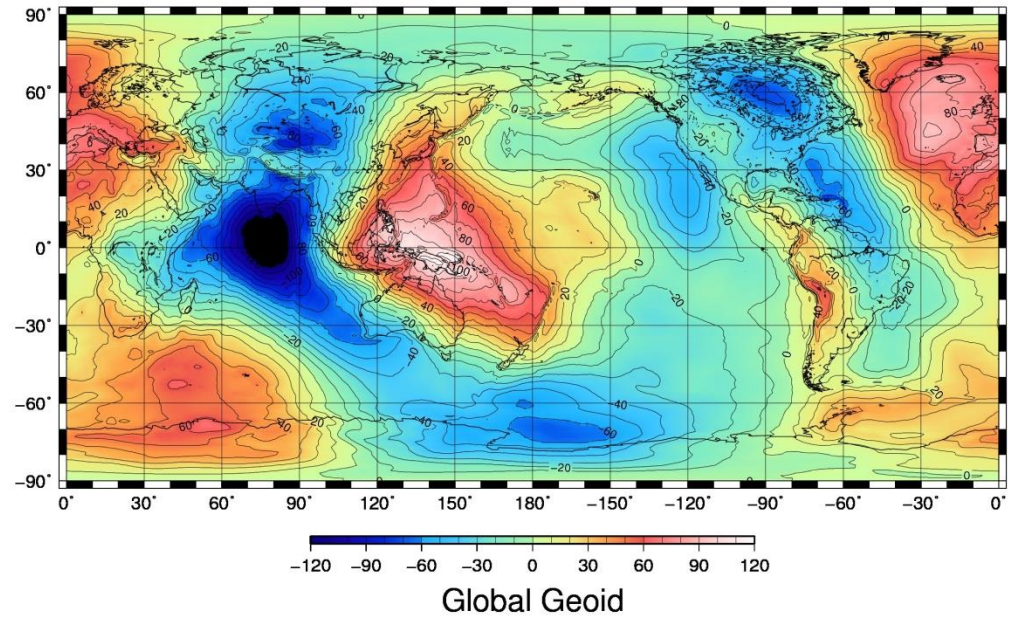
SSH

MSS_DNSC08_2M



GEOID

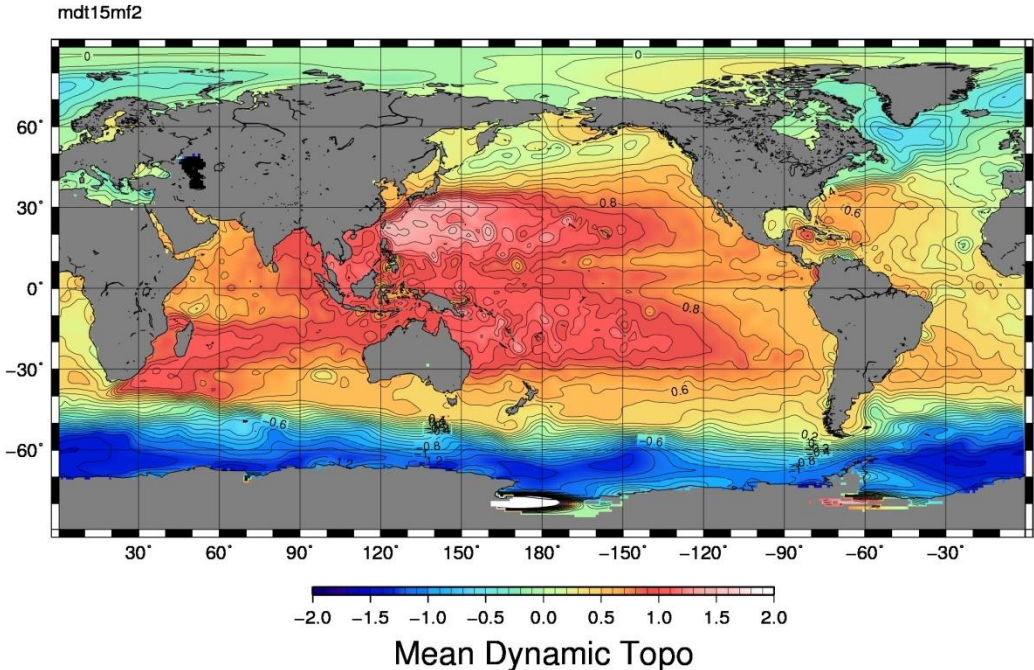
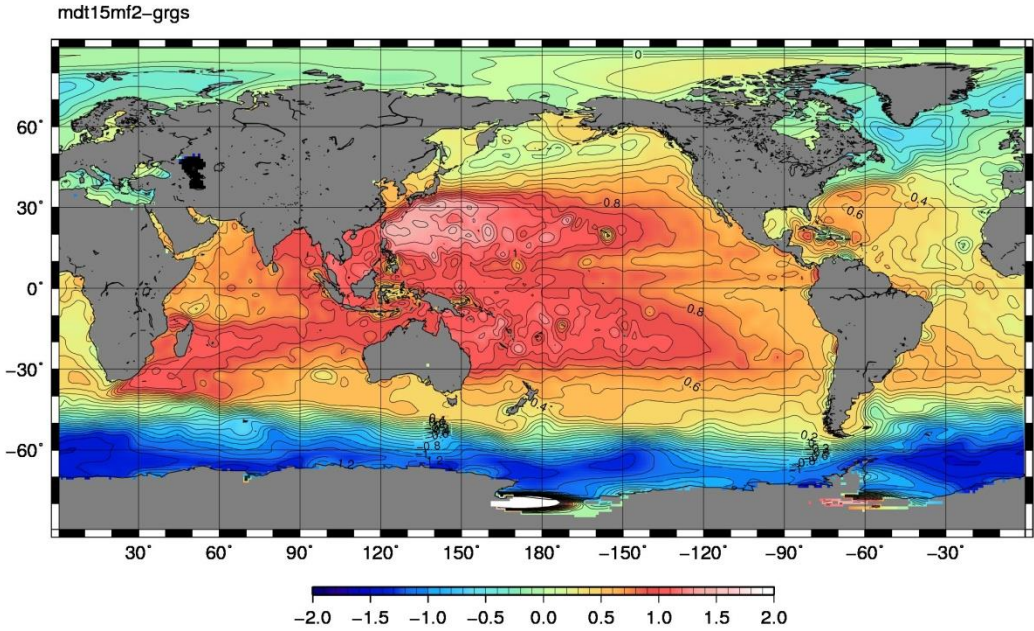
GOCE-Geoid-Direct

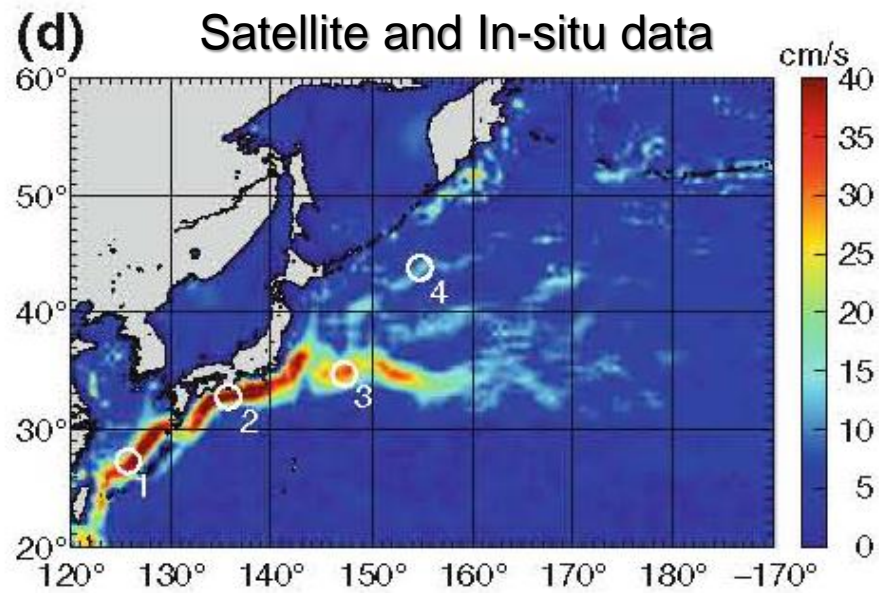
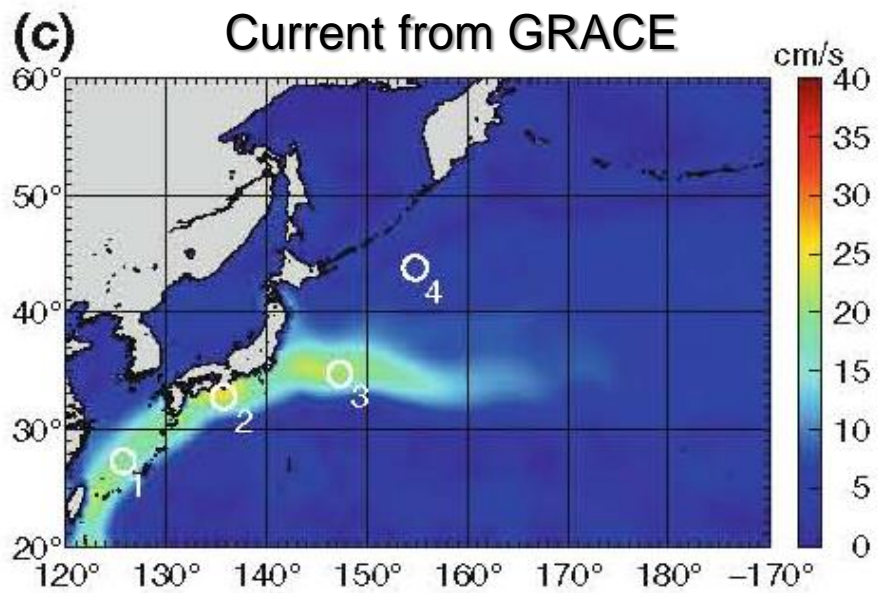
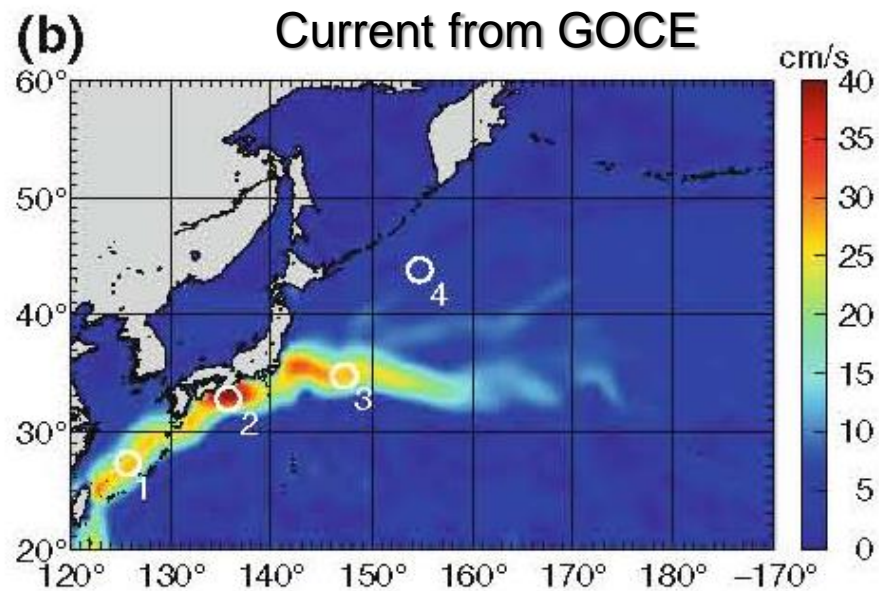
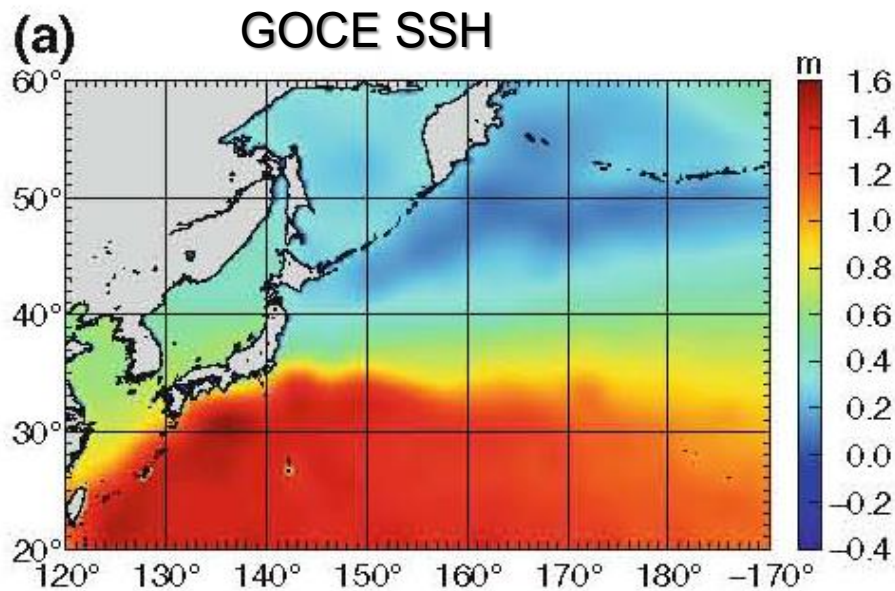


SSDT

GRACE

GOCE





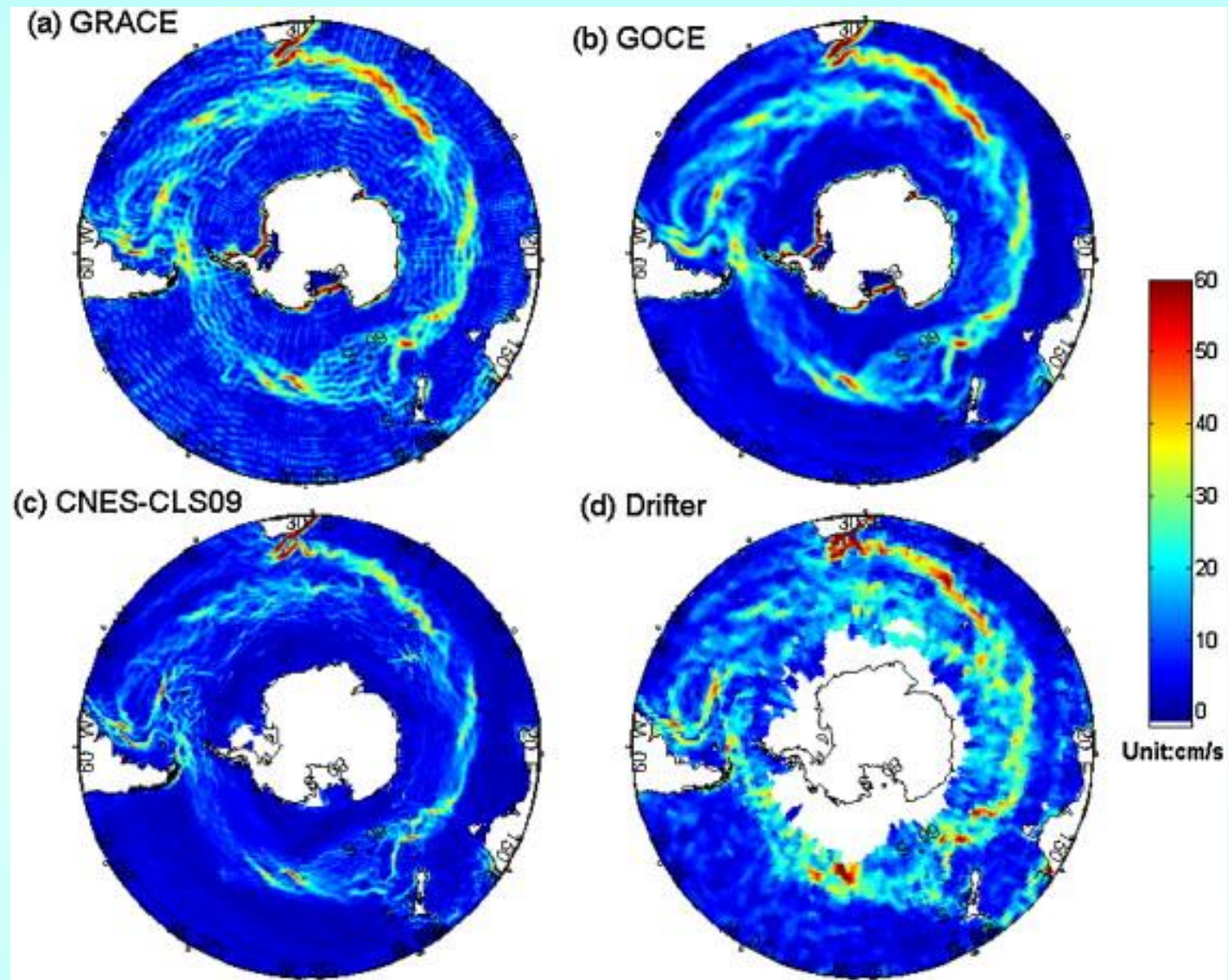


Fig. 5. Magnitude of the geostrophic velocities in the Antarctic circumpolar area. The velocity estimates from GRACE-derived geoid (ITG-Grace2010) (a), GOCE-derived geoid (GOCE-TIM3) (b), the CNES-CLS09 model (c) and the in situ drifters' measurements (d)

(Feng et al., J. Geodyn., 2013)

Global Gravity Field Models

<http://icgem.gfz-potsdam.de/ICGEM/ICGEM.htm>

Model	Year	Degree	Data	Reference	download
AIUB-GRACE02S	2009	150	S(Grace)	Jäggi et al, 2009	X
GGM03C	2009	360	S(Grace),G,A	Tapley et al, 2007	X
GGM03S	2008	180	S(Grace)	Tapley et al, 2007	X
AIUB-GRACE01S	2008	120	S(Grace)	Jäggi et al, 2008	X
EIGEN-5S	2008	150	S(Grace,Lageos)	Förste et al, 2008	X
EIGEN-5C	2008	360	S(Grace,Lageos),G,A	Förste et al, 2008	X
EGM2008	2008	2190	S(Grace),G,A	Pavlis et al, 2008	X
ITG-Grace03	2007	180	S(Grace)	Mayer-Gürr et al, 2007	X
AIUB-CHAMP01S	2007	90	S(Champ)	Prange, L. et al, 2007	X
ITG-Grace02s	2006	170	S(Grace)	Mayer-Gürr et al, 2006	X
EIGEN-GL04S1	2006	150	S(Grace,Lageos)	Förste et al, 2006	X
EIGEN-GL04C	2006	360	S(Grace,Lageos),G,A	Förste et al, 2006	X
EIGEN-CG03C	2005	360	S(Champ,Grace),G,A	Förste et al, 2005c	X
GGM02C	2004	200	S(Grace),G,A	UTEX CSR, 2004	X
GGM02S	2004	160	S(Grace)	UTEX CSR, 2004	X
EIGEN-CG01C	2004	360	S(Champ,Grace),G,A	Reigber et al, 2006	X
EIGEN-CHAMP03S	2004	140	S(Champ)	Reigber et al, 2005b	X
EIGEN-GRACE02S	2004	150	S(Grace)	Reigber et al, 2005a	X
TUM-2S	2004	70	S(Champ)	Wermuth et al., 2004	X
DEOS_CHAMP-01C	2004	70	S(Champ)	Ditmar et al, 2006	X

Model	Year	Degree	Data	Reference	download
JYY_GOCE02S	2013	230	S(Goce)	Yi et al, 2013	◆zip◆
GOGRA02S	2013	230	S(Goce,Grace)	Yi et al, 2013	◆zip◆
ULux_CHAMP2013s	2013	120	S(Champ)	Weigelt et al, 2013	◆zip◆
ITG-Goce02	2013	240	S(Goce)	Schall et al, 2013	◆zip◆
GO_CONS_GCF_2_TIM_R4	2013	250	S(Goce)	Pail et al, 2011	◆zip◆
GO_CONS_GCF_2_DIR_R4	2013	260	S(Goce,Grace,Lageos)	Bruinsma et al, 2013	◆zip◆
EIGEN-6C2	2012	1949	S(Goce,Grace,Lageos),G,A	Förste et al, 2012	◆zip◆
DGM-1S	2012	250	S(Goce,Grace)	Hashemi Farahani, et al. 2012	◆zip◆
GOCO03S	2012	250	S(Goce,Grace,...)	Mayer-Gürr, et al. 2012	◆zip◆
GO_CONS_GCF_2_DIR_R3	2011	240	S(Goce,Grace,Lageos)	Bruinsma et al, 2010	◆zip◆
GO_CONS_GCF_2_TIM_R3	2011	250	S(Goce)	Pail et al, 2011	◆zip◆
GIF48	2011	360	S(Grace),G,A	Ries te al, 2011	◆zip◆
EIGEN-6C	2011	1420	S(Goce,Grace,Lageos),G,A	Förste et al, 2011	◆zip◆
EIGEN-6S	2011	240	S(Goce,Grace,Lageos)	Förste et al, 2011	◆zip◆
GOCO02S	2011	250	S(Goce,Grace,...)	Goiginger et al, 2011	◆zip◆
AIUB-GRACE03S	2011	160	S(Grace)	Jäggi et al, 2011	◆zip◆
GO_CONS_GCF_2_DIR_R2	2011	240	S(Goce)	Bruinsma et al, 2010	◆zip◆
GO_CONS_GCF_2_TIM_R2	2011	250	S(Goce)	Pail et al, 2011	◆zip◆
GO_CONS_GCF_2_SPW_R2	2011	240	S(Goce)	Migliaccio et al, 2011	◆zip◆
GO_CONS_GCF_2_DIR_R1	2010	240	S(Goce)	Bruinsma et al, 2010	◆zip◆
GO_CONS_GCF_2_TIM_R1	2010	224	S(Goce)	Pail et al, 2010a	◆zip◆
GO_CONS_GCF_2_SPW_R1	2010	210	S(Goce)	Migliaccio et al, 2010	◆zip◆
GOCO01S	2010	224	S(Goce,Grace)	Pail et al, 2010b	◆zip◆
EIGEN-51C	2010	359	S(Grace,Champ),G,A	Bruinsma et al, 2010	◆zip◆

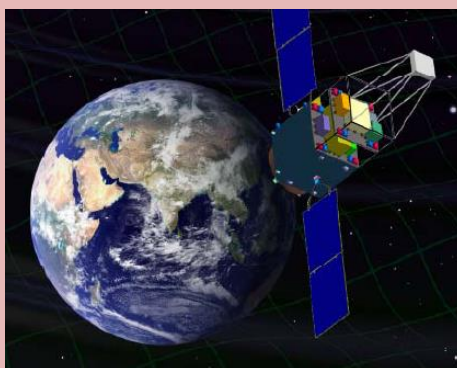
地球重力観測衛星としてのDPFのスペック

～ GOCE, GRACEとの比較



GRACE

手法	2機の衛星間の距離を測定
基線長	220km
感度	$1 \mu\text{m/s/Hz}^{1/2}$
高度	500km



DPF

手法	1軸の重力偏差計
基線長	0.3m
感度	$3.1 \times 10^{-15} \text{ /s}^2/\text{Hz}^{1/2}$
高度	500km



GOCE

手法	3軸の重力偏差計
基線長	0.5m
感度	$5 \times 10^{-12} \text{ /s}^2/\text{Hz}^{1/2}$
高度	270km

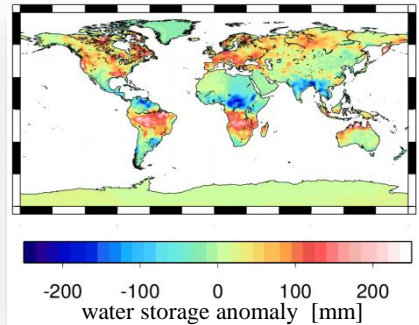
NUMERICAL SIMULATION

	Measurement error	AOD model error
Case 1	Considered	Considered
Case 2	Considered	Ignored
Case 3	Ignored	Considered

3つのケーススタディ

目的： DPFデータの陸水モニタリングへの応用

行程： 90次までの重力場を1カ月間隔で推定



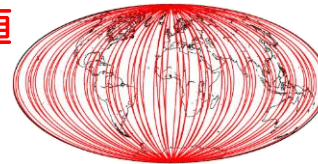
真の重力場

GLDAS陸水モデルから計算される時間変化する重力場

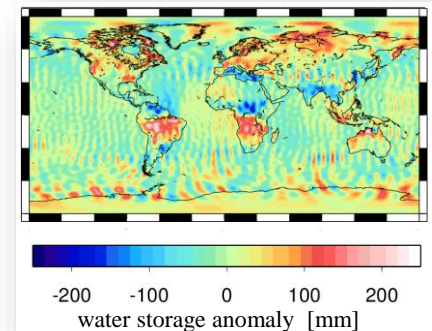


衛星軌道の**真値**

forward 計算



重力場の**推定値**



重力偏差・衛星の位置の**真値**
(5秒間隔でサンプリング)

計測誤差

- 重力偏差
- 衛星姿勢
- 衛星位置

大気海洋シグナルの補正誤差

- 大気：ECMWF - NCEP
- 海洋：OMCT - MOG2D

重力偏差・衛星の位置の計測値 (誤差を含む)

Filtering
Inversion 計算



NUMERICAL SIMULATION

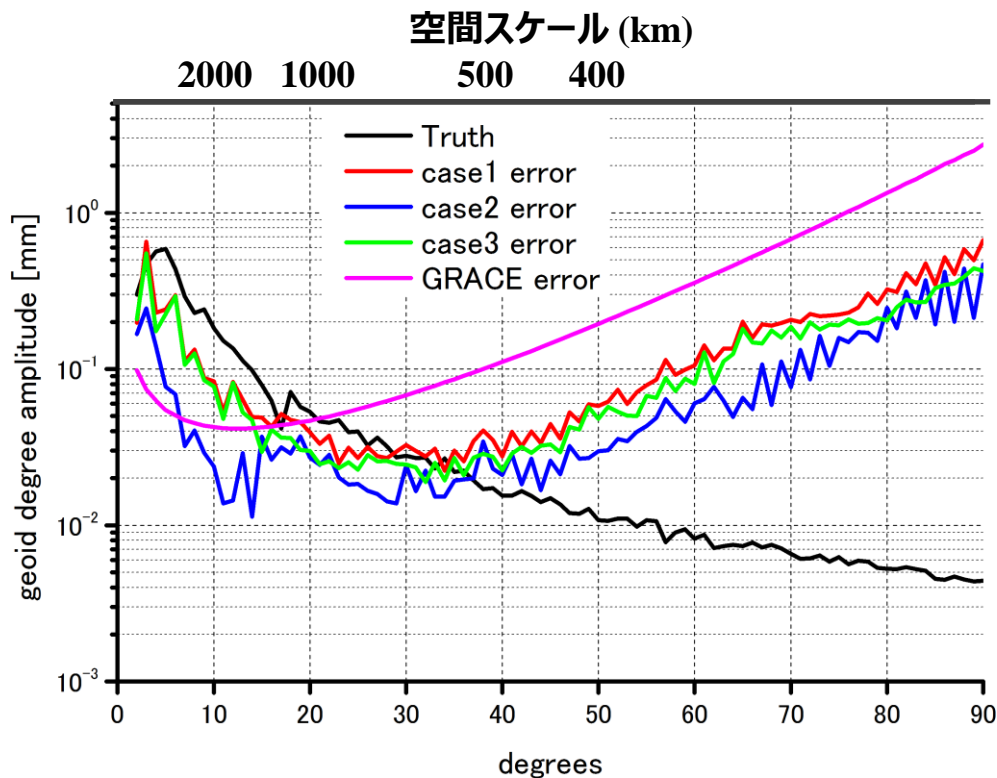
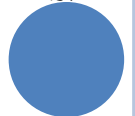


Fig. 入力した陸水シグナル、DPF観測から得られた重力場の誤差、GRACEのformal error のパースペクトル (geoid degree amplitude)

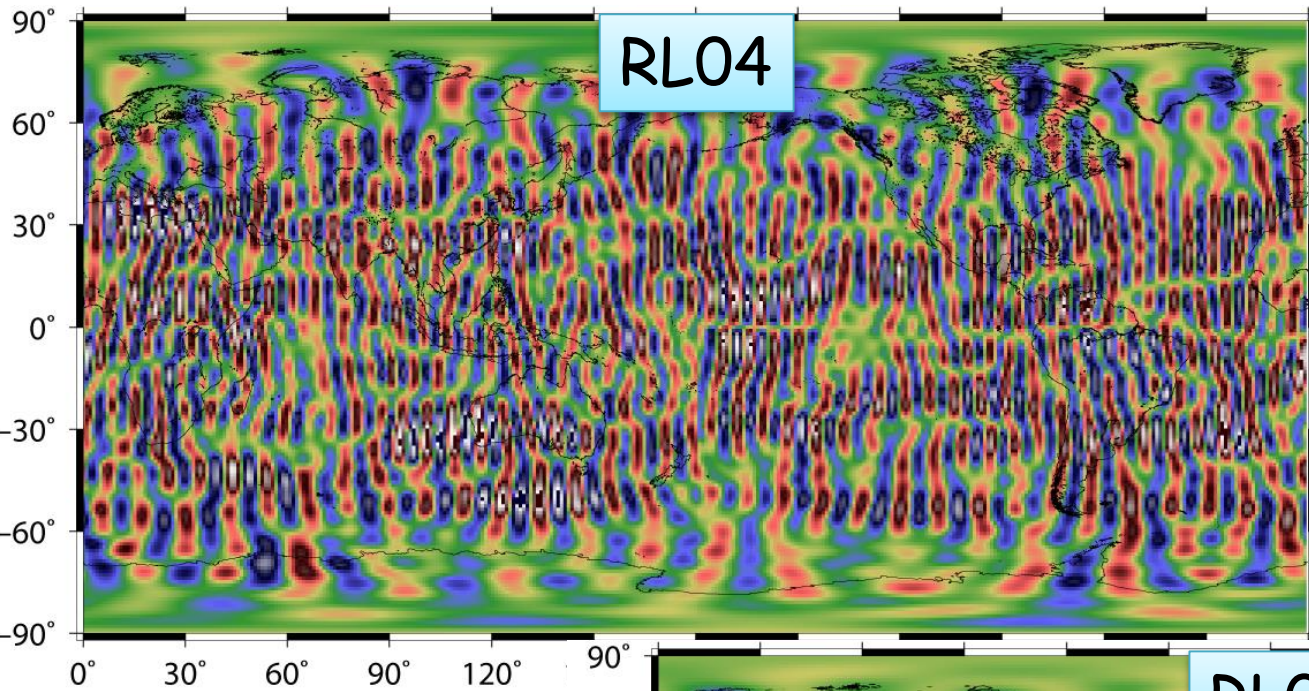
	Measurement error	AOD model error
Case 1	Considered	Considered
Case 2	Considered	Ignored
Case 3	Ignored	Considered

- 1000km以上の空間スケールでは、DPFから得られる重力場の誤差は、GRACE解の誤差より **1桁程度小さい**
- 陸水シグナルに対するSN比>1となる次数はGRACEの20次 (1000km) に対し、DPFは30次 (~660km) まで上昇
- AODモデルの誤差がDPF重力観測の主要な誤差ソース
→ **低周波帯域の計測誤差に少し余裕がある**
- 10次以下でDPF解の誤差がGRACE formal errorより大きいのはGRACE formal errorが実際のGRACE解の誤差を過小評価しているためと思われる

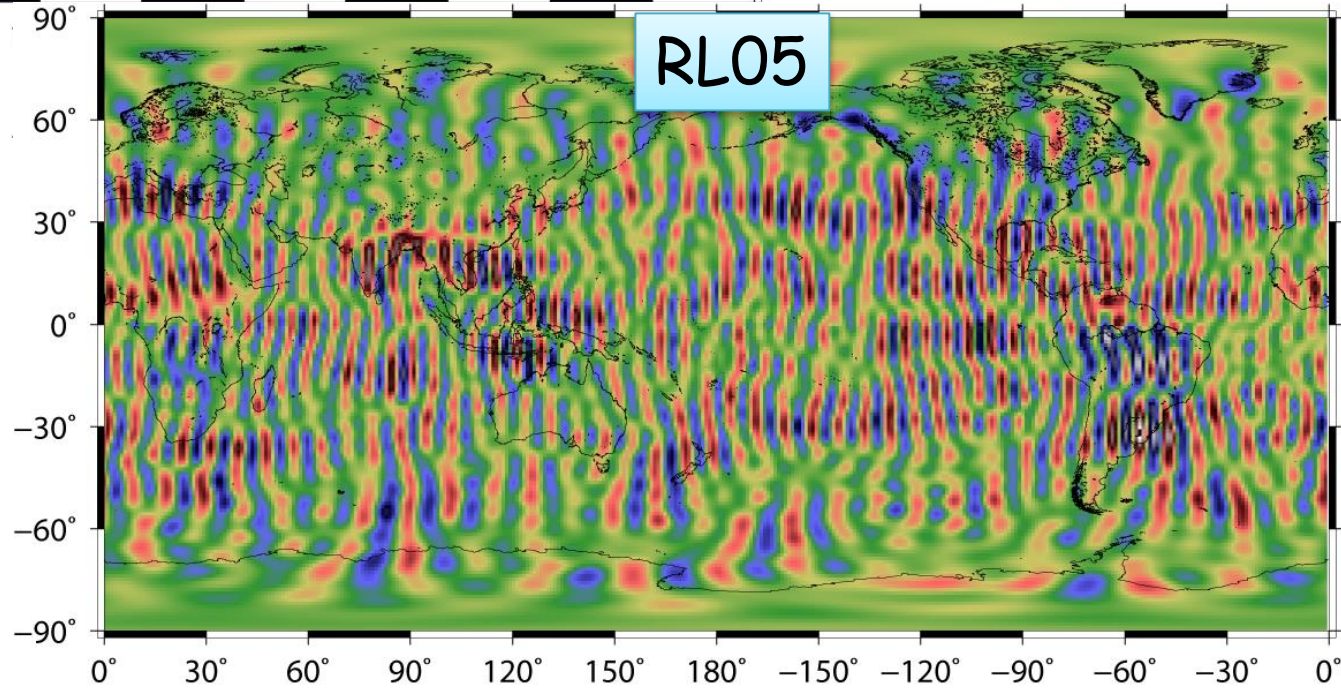


GRACE再解析

2007年6月
の表層質量変化

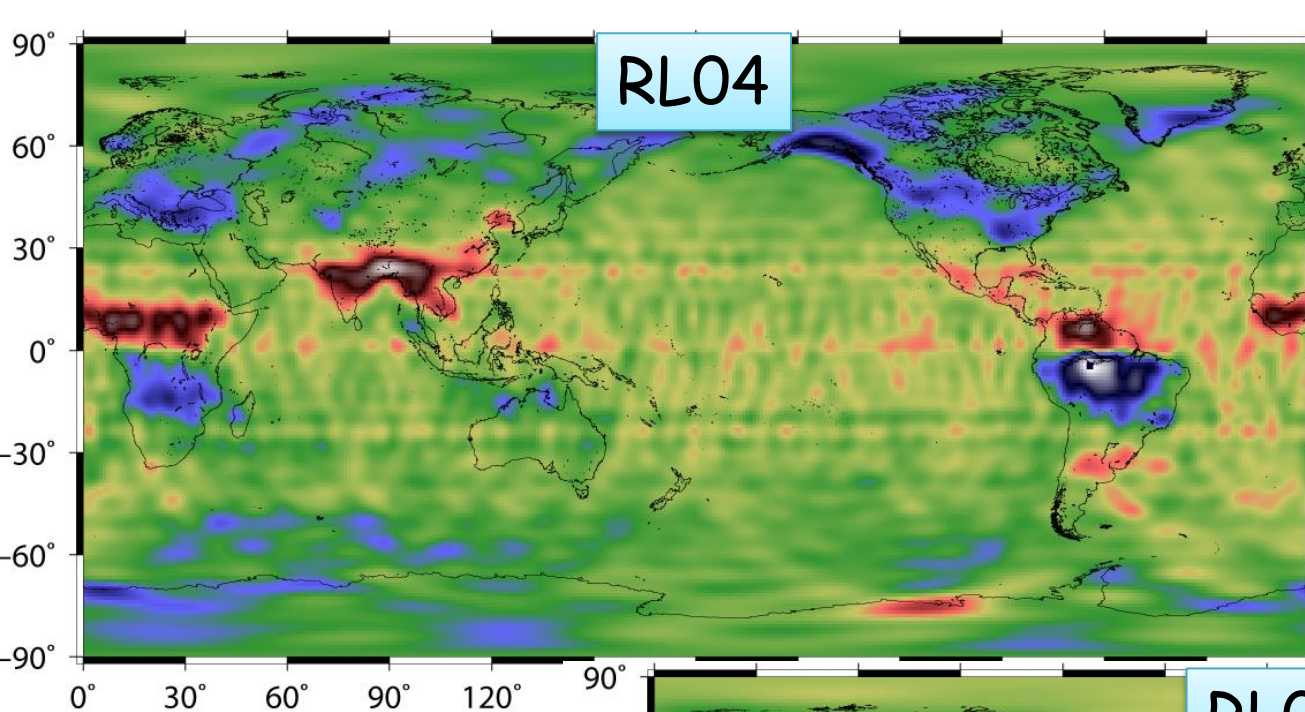


Water thickness change



Water thickness change

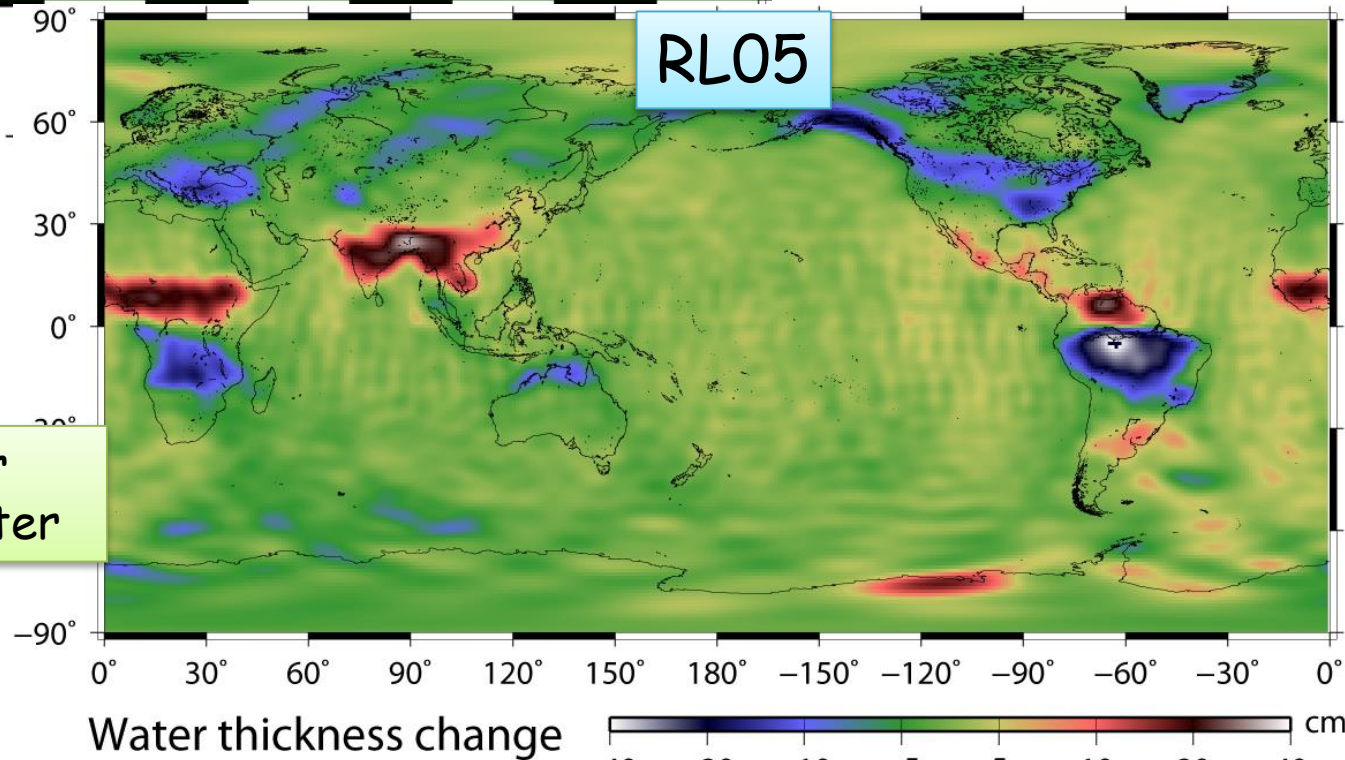




2007年6月
の表層質量変化

Water thickness change

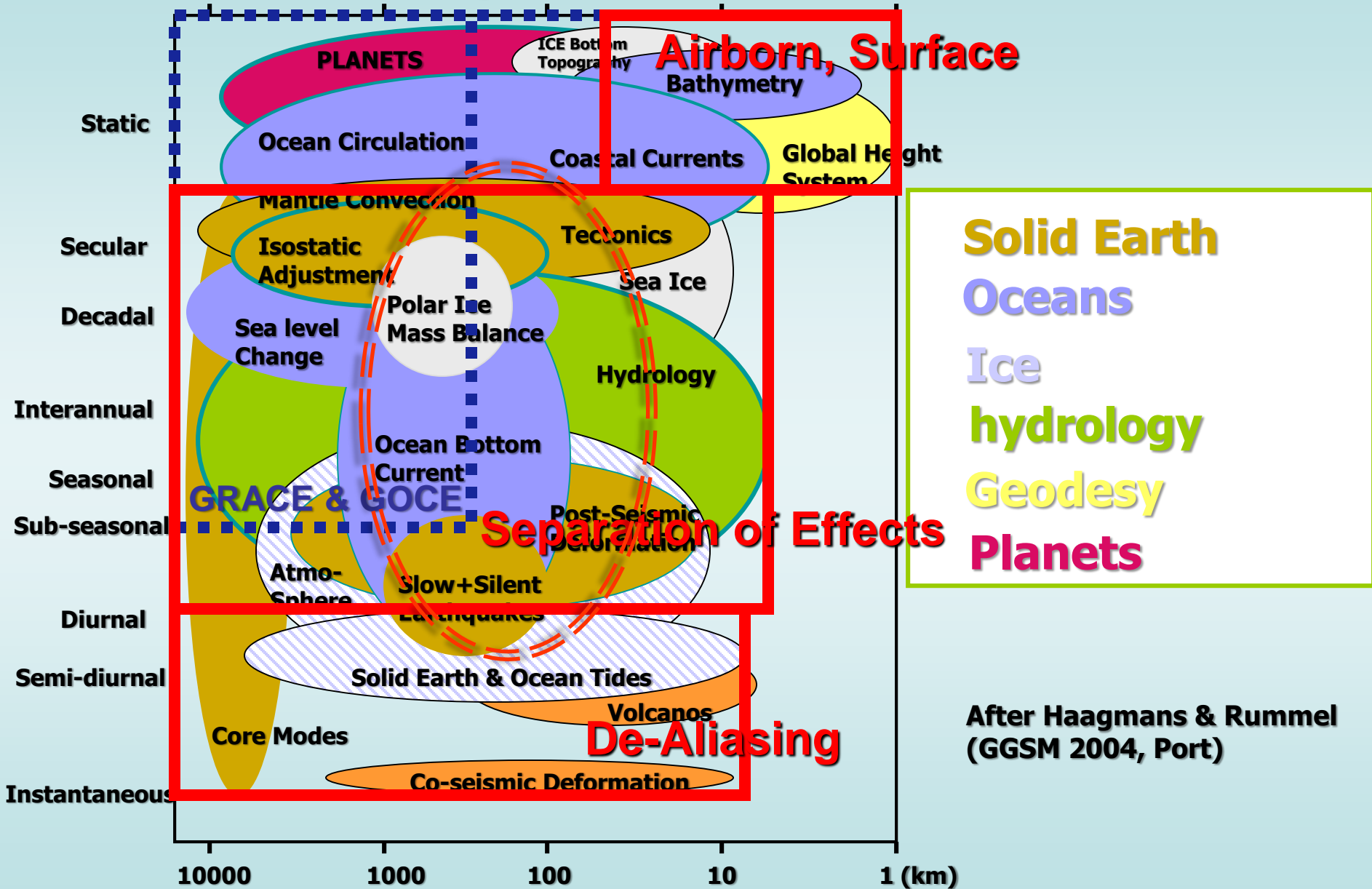
250 km Gaussian filter
+ P5M6 d-estripping filter



Gravity, Altimeter Missions

Mission	Launched	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014	2015	2016	2017	2018	Period/Until	altitude (km)	incl (deg)	
CHAMP	2000.7.15																				2010.9.19	454	87.18	
GRACE	2002.3.17																				2017??	485	89	
GOCE	2009.3.17																				2013.10-11?	250	96.5	
Swarm	2013.11.14																				4 years			
GRACE-FO	2017.12																				5 years			
TOPEX/Poseidon	1992.8.10																				2006.1.5	1336	66	
JASON-1	2001.12.7																				2013.7.2			
JASON-2	2008.6.20																							
JASON-3	7/7																							
ERS-1	1991.7.17																				2000.3.10	780	98.5	
ERS-2	1995.4.21																				2011.9.5			
Envisat	2002.3.1																				2012.4.8	790	98.6	
CryoSat	2010.4.8																				3.5 years	717	92	
Sentinel-3	2014.4																				7.25 years	815	98.65	
ICESat	2003.1.13																				2010.8.14	590	94	
ICESat-2	2016.7																				5 years			
GEOSAT																								
GEOSAT-FO	1998.02																				2008.11	800	108	
HY-2A	2011.8.15																						971	99.3
SARAL Alti-Ka	2013.02.25																				3yrs (ARGOS 5	790	98.6	

Challenges for Future Missions





Development, Operation and Analysis of Gravity Field Satellite Missions

GRACE-FO

[Science Data System](#)

[News](#)

[Literature](#)

[GRACE](#)

[Gravity Missions Studies](#)

[Satellite Payload Development and Integration](#)

[Satellite Receiving Station Ny-Ålesund](#)

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[Global Geomonitoring and Gravity Field](#) | [Topics](#)

[Development, Operation and Analysis of Gravity Field Satellite Missions](#) | [GRACE-FO](#)

Gravity Recovery and Climate Experiment-Follow-On (GRACE-FO) Mission

The Gravity Recovery and Climate Experiment-Follow-On (GRACE-FO) Mission, due for launch in August 2017, is a NASA directed mission to continue the goals of the original GRACE mission and provide continuity for the GRACE data set.

The primary goal of the GRACE-FO Mission is to obtain accurate global and high-resolution models for both the static and the time variable components of the Earth's gravity field. As in the original GRACE mission, this goal is achieved by making accurate measurements of the inter-satellite range between two co-planar, low altitude polar orbiting twin satellites using a K/Ka-Band microwave tracking system. Additionally, each satellite carries geodetic quality Global Navigation Satellite System (GNSS) receivers, a Laser Retro-Reflector (LRR) for ground station ranging, and high accuracy accelerometers to precisely measure the non-gravitational accelerations acting on the satellite.

As a secondary goal, GRACE-FO will carry a Laser Ranging Interferometer (LRI) as a technology demonstration. It will provide laser interferometry measurements of inter-satellite range changes in orbit to demonstrate laser-ranging technology in support of future GRACE-like missions. Another secondary objective is the continuation of GRACE radio occultation measurements.

The GRACE-FO project will be executed in the US under the direction of the NASA Earth Science Division (ESD) within the NASA Science Mission Directorate (SMD) and the Earth Systematic Missions Program Office at Goddard Space Flight Center (GSFC). The Jet Propulsion Laboratory (JPL) is assigned responsibility for the GRACE-FO project.

The GRACE-FO mission has significant German participation, managed by the German

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GRACE-FO Mission

- GRACE-FO Mission, due for launch in August 2017, is a NASA directed mission **to continue the goals of the original GRACE mission and provide continuity for the GRACE data set.**
- The primary goal of the GRACE-FO Mission is **to obtain accurate global and high-resolution models for both the static and the time variable components** of the Earth's gravity field.
- As a secondary goal, GRACE-FO will carry a **Laser Ranging Interferometer (LRI) as a technology demonstration.**

将来ミッション(まとめ)

- 時間変化モニタリング(氷床、陸水)
 - GRACE-FO (2017~)
 - GRACE-2 (2020以降)
 - Formation Flight Missions
- 静的な重力場
 - 空間分解能
 - 超低高度衛星
 - GOCEを超えるのは難しい
 - Cold Atom Interferometry ?

